

## **BINDING OF SO<sub>2</sub> BY SYNTHETIC SUBSTITUTED APATITES**

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### **Abstract**

The reaction of SO<sub>2</sub> with synthetic apatites was studied by TG, XRD and IR analyses at 400–1000°C. Due to an interaction of apatite with SO<sub>2</sub>, destruction of apatite and formation of CaSO<sub>4</sub> and diphosphate up to 750°C takes place. The further calcination leads to the formation of β-Ca<sub>3</sub>(PO<sub>4</sub>)<sub>2</sub> and a part of the SO<sub>2</sub> bound is lost again. The amount of SO<sub>2</sub> bound with apatite at calcination depends on the substitutions (F<sup>-</sup> ↔ OH<sup>-</sup>, PO<sub>4</sub><sup>3-</sup> ↔ CO<sub>3</sub><sup>2-</sup>, Ca<sup>2+</sup> ↔ Mg<sup>2+</sup>) in its structure.

**Keywords:** IR, sulphur oxide binding, synthetic substituted apatite, thermal analysis, XRD

### **Introduction**

Calcination of phosphate rocks has been suggested for enrichment of the ore. During calcination of phosphate rocks sulphur containing admixtures (pyrite, organic matter etc.) decompose and gaseous sulphur and sulphur oxides evolve beside the decomposition of the accompanying carbonates (calcite, dolomite) [1, 2].

The main mineral of phosphate rocks – apatite (Ap) with a general formula Ca<sub>10</sub>(PO<sub>4</sub>)<sub>6</sub>F<sub>2</sub> – may have several substitutions in the structure. The most common substitutions in the natural Aps are Mg<sup>2+</sup> and Na<sup>+</sup> for Ca<sup>2+</sup>, CO<sub>3</sub><sup>2-</sup> for PO<sub>4</sub><sup>3-</sup> and OH<sup>-</sup> for F<sup>-</sup> [3]. The reactivity of Ap depends on the substitutions in its structure. The main substituents that make the structure weaker are carbonate and magnesium. Fluorine, on the contrary, increases the stability of the structure and decreases the reactivity of Ap [3]. Formation of sulphate as well as sulphide-containing Aps was established by synthesis [3–6]. The sulphide in the Ap structure oxidises at 600–700°C and formation of Ca<sub>10</sub>(PO<sub>4</sub>)<sub>6</sub>SO<sub>4</sub> takes place. The sulphate-substituted Ap is also very unstable and decomposes above 700°C [5]. Interactions of Estonian phosphorites with the products of thermal decomposition of pyrite have been studied and introduction of sulphur into the Ap structure was assumed [7].

The aim of our study was to examine the influence of Mg<sup>2+</sup>, CO<sub>3</sub><sup>2-</sup> and F<sup>-</sup> substitutions in the apatite structure on the binding of sulphur dioxide with apatite on calcination.

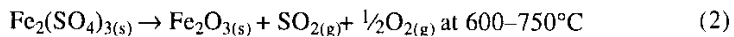
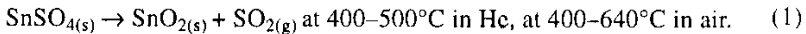
### **Materials**

The Aps were synthesized by the wet method described in [8] and identified by IR and XRD analyses as B-type carbonateapatites (CO<sub>3</sub><sup>2-</sup> substituted for PO<sub>4</sub><sup>3-</sup>). The

chemical composition of the samples can be expressed with the formulas given in Table 1. The Aps differ in the substitution in the cation site as well as in  $\text{CO}_3/\text{PO}_4$  and F/OH mole ratio.

Thermal transformations in the precipitated Aps were studied by FTIR-TG and thermogaschromatography analysis [8, 9].  $\text{H}_2\text{O}$  is released stepwise below  $550^\circ\text{C}$ . The amount of  $\text{CO}_2$  lost below  $500^\circ\text{C}$  is about 0.5% from the Ap mass. The temperature of  $\text{CO}_2$  evolution above  $500^\circ\text{C}$  depends strongly on the content of  $\text{Mg}^{2+}$  and  $\text{CO}_3^{2-}$  in the sample. The characteristic temperatures of the evolution of  $\text{H}_2\text{O}$  and of the main amount of  $\text{CO}_2$  are given in Table 1.

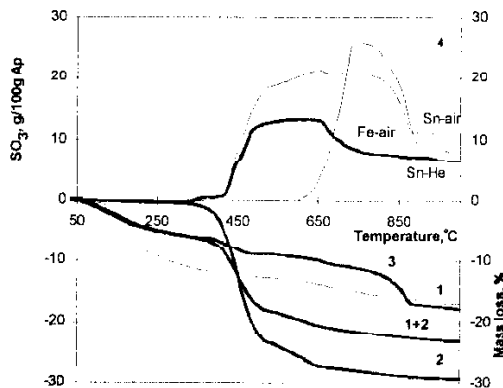
As a source of sulphur oxide,  $\text{SnSO}_4$  and  $\text{Fe}_2(\text{SO}_4)_3$  of analytical grade were used. Decomposition of the sulphates occurs by the following reactions:



## Experiment

The TG method used in the study of sulphur dioxide binding with an internal source of  $\text{SO}_2$  was described in [10]. Three runs under the same conditions were performed: 1. apatite alone; 2. sulphate alone and 3. the two together. The amount of  $\text{SO}_2$  bound is described approximately by curve 4 ( $\Delta m$ ) that exhibits the difference between the summarised mass loss of Ap and sulphate (1+2), respectively, to their mass ratios in the mixture) and the real mass loss from the third run (Fig. 1).

As at temperatures above  $650^\circ\text{C}$   $\text{SO}_3^{2-}$  is completely oxidized to  $\text{SO}_4^{2-}$  in air by the reaction (3), the calculations are performed with a presumption that  $\text{SO}_3$  was bound.

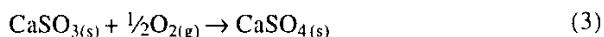


**Fig. 1** Thermogravimetric curves: 1 – apatite EM1; 2 –  $\text{SnSO}_4$  in air; 3 –  $\text{SnSO}_4 + (\text{Al}_2\text{O}_3) + \text{EM1}$  in layers in air; 1+2 – calculated curve (decomposition without interaction); 4 –  $\Delta m$  – calculated difference in masses of curves 3 and 1+2

Table I Chemical composition and thermal characteristics of the apatites

Sample	Formula	H <sub>2</sub> O evolution		CO <sub>2</sub> evolution	
		T <sub>end</sub> /°C	T <sub>start</sub> /°C	T <sub>start</sub> /°C	character
A0	Ca <sub>9,81</sub> □ <sub>1,19</sub> (PO <sub>4</sub> ) <sub>8,98</sub> (CO <sub>3</sub> ) <sub>4,02</sub> (OH) <sub>1,640</sub> □ <sub>0,36</sub>				
C1	Ca <sub>9,65</sub> □ <sub>0,35</sub> (PO <sub>4</sub> ) <sub>8,29</sub> (CO <sub>3</sub> ) <sub>0,71</sub> (OH) <sub>2,2</sub> ·2H <sub>2</sub> O	450	570	570	continuous
C2	Ca <sub>9,86</sub> □ <sub>0,14</sub> (PO <sub>4</sub> ) <sub>8,72</sub> (CO <sub>3</sub> ) <sub>0,28</sub> F <sub>0,50</sub> (OH) <sub>1,50</sub> ·1.71H <sub>2</sub> O	550	570	570	continuous
C3	Ca <sub>9,86</sub> □ <sub>0,14</sub> (PO <sub>4</sub> ) <sub>8,72</sub> (CO <sub>3</sub> ) <sub>0,28</sub> F <sub>0,89</sub> (OH) <sub>1,11</sub> ·1.53H <sub>2</sub> O	550	700	700	continuous
C4	Ca <sub>9,46</sub> □ <sub>0,54</sub> (PO <sub>4</sub> ) <sub>8,37</sub> (CO <sub>3</sub> ) <sub>0,63</sub> F <sub>1,56</sub> □ <sub>0,44</sub> ·1.70H <sub>2</sub> O	520	520	520	max at 820
E4	Ca <sub>9,33</sub> □ <sub>0,67</sub> (PO <sub>4</sub> ) <sub>8,85</sub> (CO <sub>3</sub> ) <sub>1,05</sub> F <sub>1,69</sub> (OH) <sub>0,31</sub> ·1.44H <sub>2</sub> O+0.34CaCC <sub>3</sub>	440	580	580	max at 757, 885
EM1	Ca <sub>7,13</sub> Mg <sub>1,33</sub> □ <sub>0,58</sub> (PO <sub>4</sub> ) <sub>2,91</sub> (HPO <sub>4</sub> ) <sub>1,98</sub> (CO <sub>3</sub> ) <sub>1,11</sub> (OH) <sub>2,67</sub> ·6.67H <sub>2</sub> O	550	450	450	max at 690
EM4	Ca <sub>7,82</sub> Mg <sub>1,62</sub> □ <sub>0,66</sub> (PO <sub>4</sub> ) <sub>4,81</sub> (CO <sub>3</sub> ) <sub>1,11</sub> F <sub>1,59</sub> (OH) <sub>0,4</sub> ·6.84H <sub>2</sub> O	450	450	450	max at 660

□ vacancies in the formula



The experiments were carried out with MOM (Hungary) and SETARAM (France) equipment at a heating rate of  $5 \text{ deg min}^{-1}$  in an air and He atmosphere. The mass ratio of the materials was  $\sim 56 \text{ mg of SO}_2$  evolved from sulphate per 100 mg of Ap (100 or 15 mg of Ap, respectively, on equipment).

IR spectra were obtained on a Carl Zeiss Jena IR 75 Spectrometer using the KBr pellets.

X-ray diffraction analysis was carried out with a DRON-4 Diffractometer using  $\text{CuK}_\alpha$  radiation at 40 kV, 20 mA. The samples were scanned in the range of  $8\text{--}60^\circ$  with a step size  $0.04^\circ$ .

## Results and discussion

### *Thermal analysis*

The curves of the mass change ( $\Delta m$ ) show that an increase in  $\Delta m$  begins with the beginning of decomposition of sulphate (Figs 1 and 2). In the experiments with  $\text{SnSO}_4$  the mass loss of CaAps (C1, C4) is higher than of Ap alone in the temperature range  $380\text{--}450^\circ\text{C}$ . From the precipitated apatites water evolves at temperatures up to  $450\text{--}550^\circ\text{C}$  (Table 1). Therefore, an assumption could be made that volatilization of water is accelerated as a result of interaction of Ap with  $\text{SO}_2$ . From  $450^\circ\text{C}$  up to the end of decomposition of  $\text{SnSO}_4$  ( $\sim 650^\circ\text{C}$ )  $\Delta m$  increases. The increase in the mass is higher in the air than in the He atmosphere as a result of oxidation of sulphites (Fig. 1). After  $650^\circ\text{C}$  slow mass loss begins, the course of which depends on the composition of the apatite.

In the experiments with  $\text{Fe}_2(\text{SO}_4)_3$  (Fig. 1)  $\Delta m$  has the maximum value at  $720\text{--}730^\circ\text{C}$ . By the data of IR analysis evolution of  $\text{CO}_2$  under the effect of  $\text{SO}_2$  was accelerated and occurs at lower temperatures. Therefore, the calculated amounts of the  $\text{SO}_3$  bound are not exact. However, since the amount of  $\text{CO}_2$  constitutes about 10% from the mass increase in the experiments with CaAps, and 20% on an average with Mg substituted Aps, the calculated amounts of the  $\text{SO}_3$  bound could be used for characterization of apatites.

The amount of  $\text{SO}_3$  bound by an Ap at  $400\text{--}600^\circ\text{C}$  is smaller than the amount bound at  $600\text{--}700^\circ\text{C}$ , and reaching from 15.8 to 27.7 g per 100 g Ap at  $750^\circ$  and from 4.0 to 12.7 g at  $1000^\circ\text{C}$  (Table 2), depending on the sample. The degree of the  $\text{SO}_3$  bound decreases with an increase in the fluorine content and increases with an increase in the carbonate content in the Ap structure at  $750^\circ\text{C}$ . At  $1000^\circ\text{C}$  fluorine has no remarkable effect on the amount of the  $\text{SO}_2$  bound but the latter increases with an increase in the  $\text{CO}_3/\text{PO}_4$  mole ratio in CaAp.

Mg substitution in the structure of Ap accelerates the sorption of  $\text{SO}_2$  up to  $750^\circ\text{C}$  and its evolution at higher temperatures similarly to the results obtained in the experiments with dolomite [11]. Promotion of the evolution of  $\text{SO}_2$  at lower temperatures could be a result of decomposition of  $\text{MgSO}_4$  above  $800\text{--}850^\circ\text{C}$ , as the decomposition of  $\text{CaSO}_4$  begins above  $1200^\circ\text{C}$ .

Table 2 The amount of  $\text{SO}_3$  bound by apatite (g/100 g Ap) in air

Sample	A0	C1	C2	C3	C4	E4	EM1	EM4
Mole ratio								
(Ca+Mg)/P	1.64	1.82	1.72	1.72	1.76	1.99	1.73	1.94
$\text{CO}_2/\text{PO}_4$	0.064	0.135	0.049	0.049	0.117	0.285	0.227	0.230
(Ca+Mg):(P+C)	1.63	1.61	1.64	1.64	1.58	1.69	1.41	1.58
Experiment								
With $\text{SnSO}_4$								
at 750°C								
$\text{SO}_2$		11.2			7.1			14.8
$\text{SO}_2/\text{CO}_2$		2.00			1.40			1.55
(Ca+Mg)/P		1.68			1.67			1.55
at 1000°C								
$\text{SO}_2$		8.4			5.7			6.7
$\text{SO}_2/\text{CO}_2$		1.5			1.1			0.7
(Ca+Mg)/P		1.72			1.69			1.65
With $\text{Fe}_2(\text{SO}_4)_3$								
at 750°C								
$\text{SO}_2$	26.7	27.7	22.0	20.5	15.8	19.0	26.0	25.3
$\text{SO}_2/\text{CO}_2$	146.9	4.9	10.1	9.4	3.1	1.7	2.7	2.8
(Ca+Mg)/P	1.31	1.48	1.45	1.47	1.56	1.75	1.41	1.62
at 1000°C								
$\text{SO}_2$	4.0	10.7	5.7	5.5	5.2	12.7	8.2	12.7
$\text{SO}_2/\text{CO}_2$	22.0	1.9	2.6	2.5	1.0	1.1	0.9	1.4
(Ca+Mg)/P	1.59	1.69	1.65	1.65	1.70	1.83	1.63	1.78

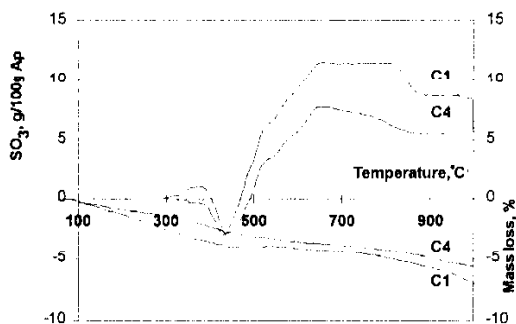


Fig. 2 TG and  $\Delta m$  curves for the samples C1 and C4 from the experiments with  $\text{SnSO}_4$

The number of  $\text{SO}_3$  moles bound exceeds the number of  $\text{CO}_2$  moles in the Ap structure (Table 2). As a result of  $\text{SO}_2$  binding and sulphate formation in the system, a part of the  $\text{Ca}^{2+}$  ions are lost from the phosphate. Therefore, in the samples with a low fluorine content ( $\text{F}/\text{OH} < 0.8$ ) the calculated mole ratio  $\text{Ca}/\text{P}$  decreases to less than 1.5 at temperatures up to  $750^\circ\text{C}$ . In  $\text{Ca}_3(\text{PO}_4)_2$  the mole ratio of  $\text{Ca}/\text{P} = 1.5$ , therefore formation of condensed phosphates is possible. Due to the partial loss of the bound  $\text{SO}_2$  on calcination up to  $1000^\circ\text{C}$ , the value of the  $\text{Ca}/\text{P}$  ratio is restored up to 1.59–1.69, but it remains lower than the value in stoichiometric Ap - 1.67. The latter indicates the possibility of formation of  $\text{Ca}_3(\text{PO}_4)_2$ .

### IR spectroscopy

In the spectra of HCAp after sorption of  $\text{SO}_2$  at  $400\text{--}600^\circ\text{C}$  in air and calcination up to  $650$  and  $750^\circ\text{C}$  all the bands characteristic of carbonate Ap are retained [3]. Besides, weak bands at  $635$  and  $670\text{ cm}^{-1}$  are observed, and the strong peak at  $1000\text{--}1100\text{ cm}^{-1}$  is wider than the Ap peak is (Fig. 3a). The  $\nu_2$  mode bands in the interval  $500\text{--}400\text{ cm}^{-1}$  are shifted and stronger. These changes in the spectra could be caused by the  $\nu_3$  and  $\nu_4$  vibrations of  $\text{SO}_4^{2-}$  group bound to Ca [12] and the disordering of the Ap structure. The spectra of the samples used for sorption of  $\text{SO}_2$  at  $600\text{--}700^\circ\text{C}$  and calcined up to  $750^\circ\text{C}$  differ in that the  $\text{CO}_3^{2-}$  bands at  $1400\text{--}1500\text{ cm}^{-1}$  are absent and the bands characteristic for  $\beta\text{-Ca}_2\text{P}_2\text{O}_7$  at  $930\text{--}1010$ ,  $720$  and  $530\text{ cm}^{-1}$  appear. At the same time the  $\text{SO}_4^{2-}$  bands are also more intensive (Fig. 3a).

In the spectra of the samples calcined up to  $1000^\circ\text{C}$  the bands of HCAp disappeared (Fig. 3b). The peaks of  $\text{SO}_4^{2-}$  and  $\beta\text{-Ca}_3(\text{PO}_4)_2$  (at  $980\text{--}1100$  and  $545\text{--}605\text{ cm}^{-1}$ ) are observed, though the relative intensities of  $\text{SO}_4^{2-}$  bands are decreased.

The changes in the spectra of HFCAp at sorption of  $\text{SO}_2$  and calcination are analogous to the changes described above. Whereas fluorine makes apatite structure more resistant, the quantity of  $\text{SO}_3$  bound is lower (Table 2) and, accordingly, the bands of  $\text{SO}_4^{2-}$  and  $\text{P}_2\text{O}_7^{2-}$  are weaker. Existence of FAp at  $1000^\circ\text{C}$  is also observed.

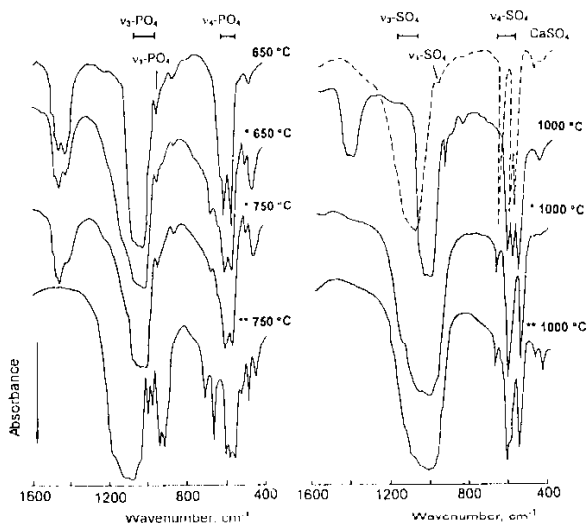


Fig. 3 IR spectra of sample C1 calcined up to given temperature: \* –  $\text{SO}_2$  bound at 400–600°C; \*\* –  $\text{SO}_2$  bound at 600–700°C.

### XRD analysis

In the samples calcined up to 750°C, after  $\text{SO}_2$  binding, inherence of  $\text{CaSO}_4$ , diphosphate ( $\text{Ca}_2\text{P}_2\text{O}_7$  or  $\text{CaMgP}_2\text{O}_7$ ) and HFAP or traces of HAP are identified. In the spectra of the samples calcined up to 1000°C the peaks of diphosphate and HAP disappear and the peaks of  $\beta\text{-Ca}_3(\text{PO}_4)_2$  [(Ca,Mg) $_3(\text{PO}_4)_2$  with Mg substituted Aps] appear. Substitution of  $\text{SO}_4^{2-}$  ions into the Ap structure in the mixture of phases was not established.

### Conclusions

Due to an interaction of apatite with  $\text{SO}_2$ , destruction of apatite, primarily of HCAPs, and formation of  $\text{CaSO}_4$  and diphosphate up to 750°C take place. The further calcination leads to reaction between diphosphate and  $\text{CaSO}_4$  and to the formation of  $\beta\text{-Ca}_3(\text{PO}_4)_2$  (of (Ca,Mg) $_3(\text{PO}_4)_2$  with Mg substituted Aps) and a part of the  $\text{SO}_2$  bound is lost again.

The amount of  $\text{SO}_3$  bound with apatite at calcination depends on the substitutions in its structure. The amount of  $\text{SO}_2$  bound is bigger when the content of  $\text{CO}_3^{2-}$  is higher and the content of F<sup>-</sup> ions in the apatite structure is lower. The content of  $\text{SO}_4^{2-}$  ions remaining at 1000°C does not depend on the fluorine content, although it is higher with a higher  $\text{CO}_3^{2-}$  content in the apatite. The evolvement of  $\text{SO}_2$  from magnesium-substituted apatites proceeds more extensively.

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